

then call the 12-month values formed of these (thus eliminating the yearly period)

$$b_1 \ b_2 \ b_3 \ . \ . \ .$$

thus:

$$b_1 = a_1 + a_2 + \dots + a_{12} = \sum_{i=1}^{12} a_i; \ b_2 = a_2 + a_3 + \dots + a_{13} = \sum_{i=2}^{13} a_i; \ . \ . \ . \ b_k = a_k + a_{k+1} + \dots + a_{k+11} = \sum_{i=k}^{k+11} a_i$$

or these values divided by 12, then the difference between two successive *b* values, or the difference of the *b* curve, will be expressed by

$$\Delta b_{i+1} = b_{i+1} - b_i = a_{i+12} - a_i$$

From this it follows that the curve of the *b* values (freed of yearly period) will rise as long as $a_{i+12} - a_i > 0$; that it will run parallel to the *x* axis as soon as $a_{i+12} - a_i = 0$; that the curve will fall as long as $a_{i+12} - a_i < 0$. If $a_{i+12} - a_i$ reaches a maximum or a minimum, this represents an epoch in the curve; in the first instance (maximum), and as viewed from the *x* axis, the curve changes from the convex form into the concave, and vice versa in the second instance (minimum).

If now, beginning with the first appearance of the period of a few years pointed out above, we can extrapolate at least for a short stretch, then we are in a position to compute in advance the corresponding monthly values, or at any rate their magnitude in relation to the value for the corresponding months of the previous year. In the space of time for which one has computed a decline of the *b* curve, the corresponding *a* value must be lower than that for 12 months previous. If the curve is passing an epoch, then the difference according to absolute values must be at a maximum.

If the *b* curve reaches its minimum, then the corresponding *a* values are like those for 12 months previous. If the *b* curve rises the *a* values are higher than 12 months previous; at an epoch this difference is greatest; at a maximum of the *b* curve the corresponding *a* values are like those 12 months previous, etc. The assumption that this is a usable method rests upon the fact that the period of a few years is so regular that it allows extrapolation to be carried out with some certainty. If one does not carry this entirely too far, but is satisfied to extend it at the most to a year (in other words, if he is content not to use all of a half period), then this method of procedure is allowable.

In Table 1, I have brought together the *a* and *b* values for the rainfall at Stockholm from January, 1921, to May, 1926. From them it is to be observed how the *b* curve oscillates: minima appear in March, 1921, October, 1922, and December, 1924; maxima in March, 1922, and November, 1923; since the beginning of 1925 the curve is rising. According to the average length of the oscillations, there should be expected a new maximum 14 months after December, 1924; that is, in February, 1926. The precipitation up to and including July, 1926, should be larger than for the preceding 12 months, and thereafter during approximately another 14 months again smaller than in the corresponding months of the previous year.

Forecasts arrived at in this manner may not infrequently fail in case one is dealing with a particular month. The values in the table show small discontinuities. It is therefore advisable not to apply the method to a shorter space of time than that of the five-month values.

Neither is it possible to fix in advance the exact time of the maxima and minima of the curve; the forecast failures that result from using this curve are serious. But in spite of this, experience has shown that the forecasts in general give good results, especially so if one content oneself with a somewhat less definite fixing of the epochs of change (at maxima and minima). In this event, in case the forecast includes a period of at least five months, useful results can be gained.

TABLE 1.—Computation of the *b* values for Stockholm, 1921-1926

Month	<i>a</i>	<i>b</i>	$\frac{\sum b_i}{5}$	Month	<i>a</i>	<i>b</i>	$\frac{\sum b_i}{5}$
1921				1923			
January	68	532	489	November	63	848	682
February	4	476	490	December	54	837	824
March	6	454	488	1924			
April	32	481	484	January	37	834	806
May	10	496	488	February	65	760	781
June	96	512	502	March	58	753	749
July	83	497	517	April	64	720	715
August	45	522	535	May	68	650	698
September	17	556	555	June	49	664	678
October	31	588	567	July	56	675	656
November	36	608	568	August	73	653	639
December	84	560	582	September	05	606	622
1922				October	40	598	598
January	53	528	606	November	23	577	601
February	29	627	622	December	38	555	602
March	40	706	645	1925			
April	64	689	685	January	48	670	599
May	30	676	663	February	43	608	609
June	48	625	646	March	11	583	626
July	51	618	625	April	56	628	623
August	144	621	599	May	45	643	607
September	96	586	580	June	27	652	639
October	14	546	583	July	171	631	642
November	23	574	577	August	11	640	637
December	33	586	579	September	70	646	636
1923				October	84	614	-----
January	46	594	591	November	38	651	-----
February	32	597	608	December	47	-----	-----
March	5	603	632	1926			
April	24	662	657	January	27	-----	-----
May	58	702	681	February	52	-----	-----
June	60	723	710	March	17	-----	-----
July	59	714	737	April	24	-----	-----
August	147	747	785	May	82	-----	-----
September	102	800	790	-----	-----	-----	-----
October	73	840	814	-----	-----	-----	-----

NOTE.—In column *b*, for example, the number shown opposite 1921, VI (June), is the sum of the known values for the months I-XII, 1921, in column *a*. In column $\frac{\sum b_i}{5}$ the number shown opposite 1921, VI, is the mean of the known values for the months IV-VIII, 1921, in column *b*.

With reference to the single-month values for Stockholm for 1921, an investigation of them, which is too long to describe here, shows that the percentage of hits is not more than 65. If instead of these monthly values we take rainfall sums for the different periods included in the forecasts, we get the result set forth in Table 2. It is evident that the forecasts were successful in ten cases out of eleven. For the sake of comparison, I have also given the normal values of precipitation for the periods concerned.

As an example of the preparation of a forecast on the basis of Table 1, we take the following case from the recent period. In February, 1925, we knew the maximum of the *b* curve in November, 1923, on the basis of which we could compute, by extrapolation, a minimum for January, 1925. Accordingly, from July, 1925, heavier rainfall than for the preceding 12 months must set in. During the months March to June, 1925, that is, precipitation must be slighter, and that from July, 1925, to February, 1926, heavier, than that which fell during the 12 months preceding. With respect to the values for single months these forecasts were verified except for August and September, 1925, and for January, 1926, as the comparison in Table 3 shows.

TABLE 2.—Results of the long-range advance computation of precipitation amounts

Period	Rainfall amounts (mm.)		
	Normal	Forecast	Actual
March, 1921–June, 1921	156	< 169	144
July, 1921–February, 1922	413	> 364	378
March, 1922–December, 1922	499	> 440	543
January, 1923–February, 1923	70	< 82	78
March, 1923–October, 1923	398	< 487	528
November, 1923–February, 1924	171	> 134	219
March, 1924–May, 1924	111	< 87	188
June, 1924–February, 1925	458	< 600	465
March, 1925–June, 1925	156	< 237	139
July, 1925–February, 1926	413	> 416	501
March, 1926–May, 1926	111	> 112	123

TABLE 3.—Example of the results of long-range advance computation of precipitation for individual months

Month	Rainfall amounts (mm.)	
	Forecast	Actual
1925		
March	< 58	11
April	< 64	56
May	< 66	45
June	< 49	27
July	< 56	171
August	> 73	11
September	> 95	70
October	> 40	85
November	> 23	38
December	> 38	47
1926		
January	> 48	27
February	> 43	52

In this case, also, the epoch falls in summer, and for it the forecast proves untrustworthy. The high July value is somewhat misleading, too, since the heavy precipitation is made up of heavy local showers. As for the period sums, Table 2 shows a good agreement, because for March–June a precipitation below 237 mm. was forecast and the actual was 139 mm.; for July, 1925, to February, 1926, the forecast gave more than 416 mm., the actual being 501 mm.

The purpose of these forecasts is not so much to give an advance computation of the precipitation for a particular month as it is to furnish an advance estimate of our lake levels for practical use in navigation and for the regulation of the lake levels. It is precisely for this purpose that the precipitation which has occurred during a long period over a larger area is quite as significant as the exact forecasting of water levels. . . . The generalized precipitation forecasts since 1915 compared with the amounts of precipitation observed in Svealand, and compared also with the corresponding previous year's rainfalls as well as with the normal values, . . . [together with] the generalized forecasts of water levels and the levels observed in Lake Siljan in Salarna² . . . show that the rainfall tendencies forecast for Svealand were verified in 13 to 14 cases out of a total of 15, and tendencies of water levels in 15 cases out of 16. Such percentages of hits may well convince one who does not demand too much of the method that it has a practical value.

² The author presents two tables embodying all these forecasts and their verifications.—B. M. V.

BLUE-SKY MEASUREMENTS

By IRVING F. HAND

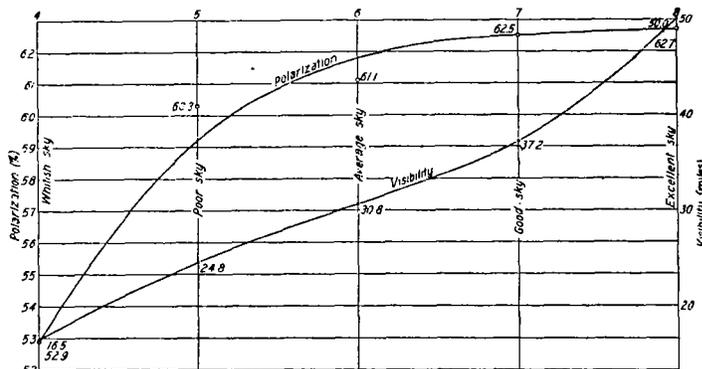
[Weather Bureau, Washington, June 30, 1927]

During the latter part of 1925 Dr. F. Linke, of the Universitäts-Institut für Meteorologie und Geophysik, Frankfurt on the Main, Germany, sent a series of color-match cards to 17 first-class meteorological services, one being the United States Weather Bureau. Commencing with January, 1926, the blueness of the sky has been noted at the Solar Observatory, American University, District of Columbia, on days when both solar radiation and skylight polarization measurements have been made. Linke states in his letter that the first measurements of the blueness of the sky of which we have record were made by H. B. de Saussure in Europe 150 years ago, and just a hundred years later the work was taken up by H. Wild. Although measurements of this kind have been made more or less frequently in Europe since the resumption of this research by Wild, but few have been made in this country.

The color-match charts consist of a series of 14 cards of graded hues ranging from almost white to a very dark blue, care having been taken to insure permanency of color and trueness of shade. The observation consists in selecting a card which when held against the blue sky 90° from the sun and in its vertical will most nearly match the color of the sky, which at this point is, generally speaking, the darkest. It is also the point of maximum polarization of skylight.

At Washington all the observations made to date cover the range of only the five cards numbered 4 to 8, inclusive. Skies whiter than that represented by card No. 4 are unfit for solar radiation measurements at

normal incidence, while skies darker than No. 8 have not been observed at this station. We may therefore, for the sake of convenience, arbitrarily designate the cards as follows: No. 4, whitish sky; No. 5, poor sky; No. 6, average sky; No. 7, good sky; and No. 8, excellent sky.



As will be seen from the figure there appears to be a very close relation between the blueness of the sky and horizontal visibility. This is to be expected as water vapor and dust both whiten the sky and decrease visibility.

The relation between sky color and the elements, visibility, polarization, solar radiation intensity at 2 air mass (zenith distance of the sun 60°), dust particles per cubic centimeter, relative humidity, vapor pressure, and